

Role of various geomorphic parameters in landslide susceptibility of the Balason river basin of Darjeeling Himalaya, India

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ABSTRACT

Occurrences of landslides are affected by geomorphic, tectonic and hydrologic parameters. Balason basin of Darjeeling Himalaya, exhibits these parameters which are analyzed in the context of landslides occurrences. In the present work, we carried out an overlay analysis of various data layers such as elevation, slope, geology, geomorphology, soil, rainfall, drainage density and drainage frequency with the landslide distribution data layer to assess the probability of class of the landslide causative factors. Frequency ratio (FR) value estimate, for both landslide affected pixels and total pixels of a class, help in establishing the relationship between the probability of landslide and each class of the landslide causative factor. The result showed that, those areas which are characterized by elevation of 2130 – 2603 m, slope of 47.00 – 71.00°, rainfall of 2559 – 2617 mm, drainage density 10.60 – 15.10 km/sq. km, 70.35 – 104.30 no. of stream/sq.km drainage frequency are registered with high frequency ratio and high landslide susceptibility. In addition, the probability of landslide occurrences is also found high in Darjeeling gneiss, lower hill and fine loamy to coarse loamy soil textural area.

Key words: Frequency ratio (FR), Susceptibility, RS and GIS, Balason river basin.

INTRODUCTION

Landslides can be defined as the movement of mass of rock, debris or earth materials down the slope under the influence of gravitational force. Many researchers in India (Nautiyal, 1966, Sengupta, 2000, Mandal and Maiti, 2014) carried out demand based studies in mountainous and hilly areas and tried to find out the causative factors and consequences of landslide occurrences and recommended some mitigation practices to reduce its hazardous impact. In Darjeeling, the spatial extents of landslides are increasing day by day, causing severe damage to lives and properties. The Balason river basin is not an exception to these phenomena, which is characterized by distinct geographical environment. The evolution of landforms is largely dependent on its geological structure, lithology, relief configuration, climatic characteristics and bio-physical processes. Researchers like Tamang (2013) and Lepcha (2012) carried out studies on geophysical set up of Balason River basin. This basin has been experiencing active denudational processes since its origin because of high relief diversity, tectonic influences, human interferences, stream activities, monsoonal rainfall and so on. Free face and rectilinear facets of slope of the landforms situated in higher altitude, especially cliffs, spurs and tea garden areas are more instable in nature of this basin. However, in this study an attempt has been made to understand the role of various geomorphic attributes i.e. elevation, slope, geology, geomorphology, soil, rainfall, drainage density and drainage frequency, to visualize the

nature and extent of variations in landforms characteristics and landslide susceptibility.

STUDY AREA

The Balason river basin, situated in the western part of the Darjeeling district, that suffered several large and small landslides and affected many properties and lives, was chosen as the study area (Figure 1). The major right bank tributary of the Mahananda river and the Balason river, originates from Lepchajagat (2361 m) which is located on the Ghum-Simana ridge. Balason river basin covers parts of Darjeeling district in West Bengal having an area of 378.45 sq. km. The main Balason river flows north to south-east, having length of about 51.92 km and joins Mahananda River at around 26°41'28'' N and 88°24'15'' E. The highest elevations are 2613 m (SE corner) and lowest 105 m, in Southern part of the basin. The area above 30° slope consists of comparatively harder gneissic rocks and the area in between 0°-5° have the permeable subsurface materials. The average rainfall of the basin is 2300 mm (range 2000-5000 mm), which is dependent on southwest monsoon and occurs mostly in the months of June to September. The mean annual temperature is about 20.94° C. The northern uplands and southern lowlands have a mean annual temperature of 12° C and 24.70° C respectively (Lama, 2003). The area is predominately represented by very shallow to deep soils. The major portion of the basin (about 74%) consists of Darjeeling gneisses, Daling series rocks and the Damuda series.

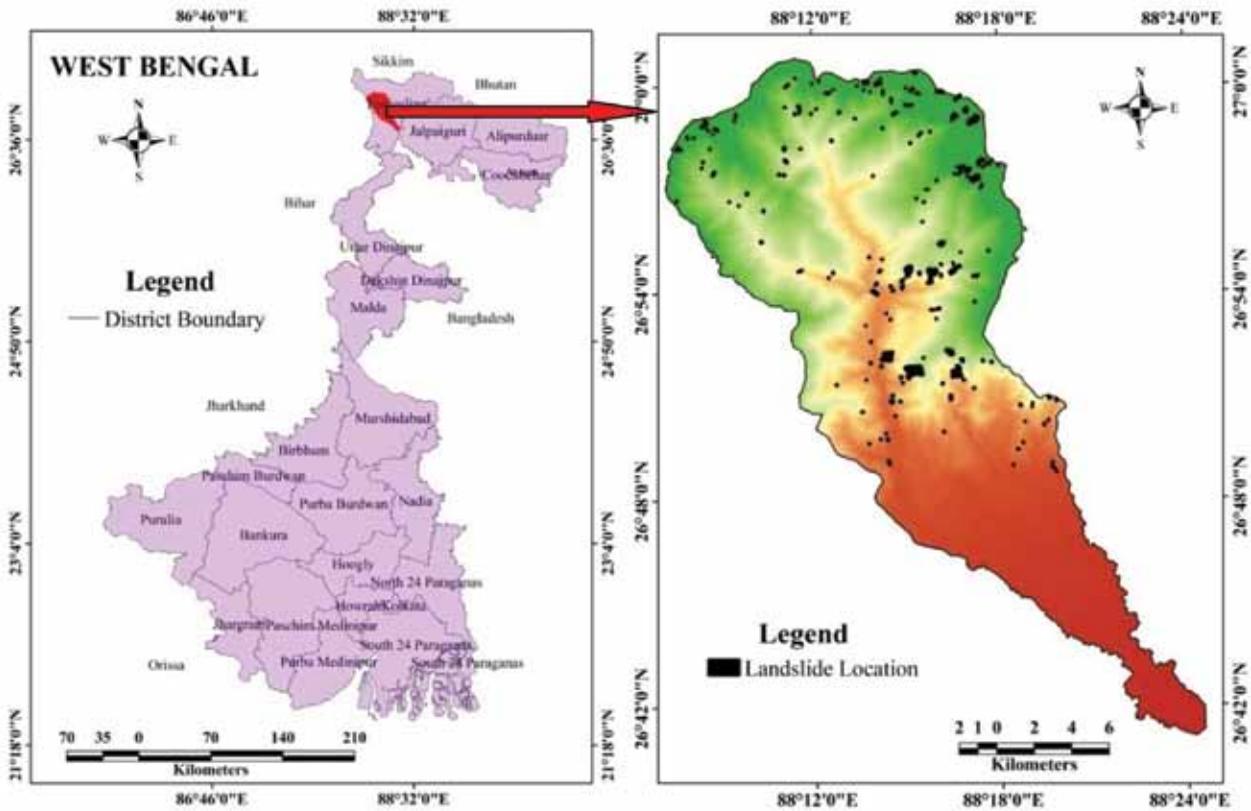


Figure 1. Location map of the study area.

METHODOLOGY

In the present work, eight parameters i.e. elevation, slope, geology, geomorphology, soil, rainfall, drainage density and drainage frequency have been used to assess the role in landslide susceptibility. Basically, landslide susceptibility refers to the possibility or chance or likelihood of potential landslides in a given area of interest based on the local terrain conditions (Grozavu et al., 2013). Elevation and slope data layers were directly derived from SRTM DEM (30m spatial resolution) in Arc GIS 10.1 environment which was downloaded using <https://earthexplorer.usgs.gov> link. The soil and geological maps were collected from NBSS (National Bureau of Soil Science and Land Use Planning, Kolkata) and GSI (Geological Survey of India, Kolkata) respectively and digitized in Arc GIS 10.1 platform. Geomorphological map was digitized from Unpublished Ph.D. Thesis of Lama (2003) using Arc GIS environment. Rainfall data was downloaded from <http://www.worldclim.org> link in Continuous raster data format and imported in Arc GIS environment. Drainage network was digitized from Survey of India (SOI), Kolkata. Map nos. 78a/4, 78a/8, 78b/1, 78b/5 and 78b/6 and was checked with and google earth image (2015). Inverse distance weighted (IDW) method was incorporated to prepare drainage density

and drainage frequency map in Arc GIS 10.1 software. The database of landslide distribution map was generated from Google earth historical imageries (2000-2016), Landsat 7 ETM+ imageries (2000-2005), Landsat 4-5 TM imageries (2006-2007, 2009-2011), LISS III imageries (2008, 2012-2013), Landsat 8 OLI (2014-2015), High resolution Sentinel-2 imagery (2016), Arc Map high resolution world imagery and intensive field survey data (2015-2016). Total 295 landslide polygons were mapped in an area of 378.45 sq.km (Table 1). The average area of the landslides was 3762.90 m².

To establish the relation between each factor and landslide occurrences, the frequency ratio (FR) value was used (Karim et al., 2011).

$$Fr_i = \frac{\{N \text{ pix } (Si)/N \text{ pix } (Ni)\} * 100}{\{\sum N \text{ pix } (Si)/\sum N \text{ pix } (Ni)\} * 100} \dots\dots\dots (1)$$

Where, N pix (Si) is the number of pixels containing landslide in each class (i), N pix (Ni) is the total number of pixels having class (i) in the whole basin, or simply the percentage of landslide pixels to total pixels in a class, $\sum N \text{ pix } (Si)$ total number of pixels in each class and $\sum N \text{ pix } (Ni)$ total number of pixels in the whole basin or simply percentage of pixels to total pixels in class. To calculate the number of landslide pixels of each class of the respective

Table 1. Area and location of the landslides in Balason river basin from 2000-2016.

S.No.	Area of the slides (Sq. m)	X coordinate of centroid of the slides area	Y coordinate of centroid of the slides area	S.No.	Area of the slides (Sq. m)	X coordinate of centroid of the slides area	Y coordinate of centroid of the slides area
1	707.199	88.2957	26.866	149	1062.05	88.2308	26.8763
2	1241.36	88.3238	26.8357	150	9596.98	88.2323	26.9005
3	31.7915	88.328	26.8172	151	50.3166	88.1552	26.9443
4	39.2974	88.3282	26.8173	152	1064.16	88.2108	26.9096
5	74.7919	88.3277	26.8153	153	1842.76	88.1723	26.9248
6	202.192	88.3303	26.8127	154	382.255	88.1405	26.9607
7	443.925	88.329	26.8144	155	395.708	88.1413	26.9636
8	182.456	88.1269	26.9448	156	244.159	88.1268	26.9694
9	76.8791	88.2955	26.9618	157	656.173	88.1278	26.9708
10	135.181	88.2954	26.9617	158	2051.94	88.129	26.9697
11	18.0313	88.2936	26.9614	159	826.188	88.1328	26.9755
12	4.79447	88.2935	26.9614	160	1431.73	88.1648	26.9482
13	115.793	88.3003	26.9606	161	2284.52	88.1736	26.9398
14	17.4485	88.2958	26.9627	162	244.237	88.1939	26.911
15	6.26611	88.2958	26.9624	163	473.227	88.2084	26.9073
16	61.9584	88.2918	26.9602	164	1118.3	88.1849	26.9571
17	21.4497	88.2922	26.9603	165	1343.13	88.1822	26.9573
18	20.2991	88.2894	26.9611	166	58.4183	88.1584	26.9833
19	28.2176	88.2895	26.9629	167	75.5655	88.159	26.9858
20	71.8128	88.2903	26.9623	168	114.422	88.1588	26.9855
21	23.3692	88.2905	26.9627	169	1045.16	88.1908	26.9865
22	15.4175	88.2908	26.963	170	72.5905	88.1933	26.9869
23	35.9775	88.291	26.9631	171	162.601	88.1935	26.9898
24	21.1504	88.2911	26.9633	172	1738.58	88.1929	26.99
25	92.1921	88.2908	26.9639	173	167.346	88.1937	26.988
26	32.9415	88.2913	26.9636	174	467.586	88.1856	26.9898
27	53.7706	88.2914	26.9641	175	530.686	88.1938	26.9698
28	17.3334	88.2918	26.9632	176	151.251	88.1858	26.9994
29	38.7774	88.2922	26.9635	177	108.052	88.187	27.0002
30	51.2226	88.2929	26.9635	178	158.193	88.1876	27.0007
31	68.9653	88.2852	26.9596	179	694.12	88.209	26.9963
32	379.161	88.2814	26.9609	180	509.409	88.2095	26.9967
33	1699.4	88.2809	26.9601	181	79.7281	88.2076	26.9966
34	2346.11	88.28	26.9597	182	51.9661	88.2119	26.9981
35	358.216	88.2803	26.9581	183	154.707	88.2041	27.0023
36	282.583	88.2857	26.9548	184	180.331	88.2003	27.0012
37	39.415	88.2862	26.9543	185	237.411	88.1997	26.9865
38	39.1453	88.2856	26.9538	186	192.888	88.2024	26.9888
39	22.5493	88.286	26.9536	187	1740.5	88.2106	26.9734
40	37.5993	88.2872	26.9537	188	337.62	88.243	26.9935
41	35.2606	88.288	26.954	189	383.982	88.2563	27.0058
42	48.6892	88.288	26.9544	190	31.9148	88.2553	27.006
43	166.631	88.2877	26.9547	191	233.374	88.2374	27.0011
44	42.7794	88.2871	26.9549	192	248.928	88.2373	27.0009
45	27.9714	88.2867	26.9549	193	65.6938	88.2373	27.0031
46	22.5377	88.2868	26.955	194	1664.16	88.2354	27.0014
47	41.6531	88.287	26.9551	195	64.8215	88.2367	27.0014
48	77.0869	88.2876	26.9552	196	2662.13	88.2354	27.0007

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S.No.	Area of the slides (Sq. m)	X coordinate of centroid of the slides area	Y coordinate of centroid of the slides area	S.No.	Area of the slides (Sq. m)	X coordinate of centroid of the slides area	Y coordinate of centroid of the slides area
49	7.44129	88.2882	26.9551	197	256.755	88.2352	27.0002
50	48.5958	88.2885	26.9552	198	910.659	88.2346	27.0004
51	76.4866	88.2883	26.9555	199	94.6291	88.2305	27.0015
52	23.7807	88.2881	26.9558	200	166.369	88.2298	27.001
53	179.457	88.2893	26.956	201	28.7871	88.2286	27.0013
54	41.638	88.2896	26.956	202	13.5587	88.2304	27.0019
55	42.7744	88.2895	26.9563	203	360.114	88.2283	27.0012
56	25.8678	88.2889	26.9542	204	73.1794	88.2286	27.001
57	80.355	88.2899	26.9543	205	176.405	88.2246	26.9997
58	11.6655	88.2898	26.9544	206	57.8003	88.2282	26.9979
59	13.3572	88.29	26.9543	207	29.0432	88.2317	26.9969
60	52.2182	88.2881	26.9541	208	645.099	88.2233	26.9841
61	17.4687	88.2869	26.9531	209	1045.4	88.2148	26.9871
62	75.9915	88.2793	26.9568	210	106.941	88.254	26.9982
63	201.801	88.2791	26.9567	211	217.025	88.258	26.9959
64	434.811	88.2768	26.957	212	299.998	88.2566	26.996
65	5056.89	88.2778	26.9845	213	32.9966	88.2548	26.9982
66	653.957	88.2758	26.9844	214	9.42264	88.2545	26.9977
67	133.234	88.2764	26.9853	215	45.1092	88.2539	26.9973
68	78.322	88.2765	26.9855	216	6.65395	88.2533	26.997
69	71.0966	88.2772	26.986	217	207.835	88.2505	26.9906
70	92284.2	88.2872	26.9825	218	67.6416	88.2512	26.9901
71	7635.69	88.2885	26.9846	219	61.1047	88.253	26.9904
72	23530.9	88.283	26.9878	220	300.908	88.2541	26.9903
73	1342.71	88.2775	26.9928	221	30.4599	88.243	26.9917
74	2953.5	88.2798	26.9944	222	133.583	88.2508	26.9986
75	20926.3	88.2692	26.9967	223	465.7	88.2459	26.9867
76	649.424	88.2385	26.9974	224	36.006	88.2444	26.9866
77	68.4146	88.2384	26.9956	225	11.4642	88.2444	26.9865
78	982.645	88.2692	26.8557	226	37.306	88.2438	26.9857
79	315.397	88.287	26.8951	227	319.488	88.2328	26.9563
80	7041.19	88.2786	26.9122	228	602.203	88.253	26.9719
81	10350.4	88.2876	26.9076	229	209.096	88.2575	26.9754
82	241.655	88.1912	27.0038	230	174.979	88.2584	26.9764
83	1791.87	88.1862	26.9635	231	675.168	88.2589	26.9768
84	1133.26	88.1853	26.9642	232	253.721	88.2636	26.9773
85	7687.55	88.1851	26.9697	233	172.194	88.2756	26.9829
86	35391.3	88.2693	26.9233	234	41.7062	88.2757	26.9832
87	8668.89	88.1469	26.9579	235	377.121	88.2381	26.9792
88	96.2151	88.2784	26.9603	236	13.6503	88.2592	26.934
89	153.184	88.2796	26.959	237	103.161	88.26	26.9435
90	327760	88.254	26.862	238	83.3302	88.256	26.9453
91	86276.3	88.2515	26.9106	239	39.2229	88.2623	26.9495
92	178297	88.2394	26.869	240	11.7083	88.2612	26.9495
93	51441.1	88.2733	26.8709	241	292.799	88.267	26.9538
94	34320.9	88.2753	26.9101	242	75.0716	88.2831	26.9577
95	26807.8	88.2629	26.9093	243	92.8112	88.284	26.9581
96	50497.7	88.265	26.9098	244	141.598	88.2848	26.9585
97	1356.49	88.1428	26.9831	245	44.7715	88.2842	26.9587
98	302.838	88.2356	26.8291	246	65.5269	88.2836	26.9595

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99	4286.21	88.2783	26.8624	247	30.6295	88.2834	26.9586
100	67.006	88.2906	26.9501	248	19.372	88.2831	26.9592
101	207.329	88.2915	26.9515	249	153.508	88.2832	26.9595
102	3608.52	88.2776	26.8639	250	47.3037	88.2833	26.9596
103	594.788	88.2946	26.9197	251	23.1429	88.2912	26.963
104	8094.87	88.28	26.8669	252	59.9813	88.2839	26.9556
105	293.191	88.2766	26.8611	253	156.932	88.2373	26.916
106	167.62	88.1591	26.9864	254	2103.85	88.243	26.9009
107	205.681	88.2295	26.827	255	1549.07	88.2619	26.9045
108	31.4893	88.1604	26.9424	256	63.8999	88.2639	26.9048
109	529.442	88.226	26.8423	257	251.299	88.2633	26.9045
110	6607.22	88.1428	26.9818	258	64.6839	88.2636	26.9064
111	187.824	88.142	26.9805	259	2342.79	88.2632	26.9066
112	209.821	88.1405	26.9805	260	1326.81	88.2689	26.908
113	81.1721	88.1412	26.9814	261	212.756	88.269	26.9133
114	101.34	88.145	26.9699	262	314.619	88.2756	26.9138
115	48.6915	88.1362	26.9754	263	43.8935	88.2797	26.919
116	25.7825	88.1465	26.9772	264	2120.05	88.2505	26.9079
117	165.093	88.1442	26.9789	265	242.618	88.2634	26.8916
118	93.9419	88.2303	26.917	266	252.802	88.2644	26.8913
119	717.515	88.1491	26.9788	267	262.503	88.2667	26.8917
120	458.494	88.1486	26.979	268	1719.26	88.2677	26.892
121	70.2184	88.2336	26.9036	269	43.2021	88.2662	26.887
122	548.314	88.1479	26.9797	270	27.6787	88.259	26.8683
123	1859.56	88.2423	26.85	271	6.49692	88.2282	26.8839
124	29211	88.2751	26.8605	272	407.037	88.2477	26.8563
125	854.153	88.2477	26.9021	273	279.718	88.2122	26.8555
126	11015	88.2418	26.8472	274	28.7389	88.2281	26.8577
127	328.532	88.1923	26.9103	275	5350.87	88.2772	26.8624
128	252.397	88.2343	26.9044	276	30543.9	88.2777	26.8597
129	3010.4	88.2363	26.8644	277	243.052	88.2809	26.8573
130	1247.26	88.2468	26.8589	278	260.571	88.2803	26.8575
131	789.487	88.2464	26.8472	279	212.472	88.2862	26.8534
132	14854.8	88.2509	26.8593	280	14374.7	88.2736	26.8491
133	1947.98	88.2402	26.8172	281	289.395	88.2698	26.8325
134	416.474	88.2414	26.8869	282	402.556	88.2435	26.8402
135	63.2228	88.2392	26.8897	283	196.266	88.3069	26.8162
136	842.363	88.2418	26.9028	284	1235.7	88.3111	26.8451
137	179.9	88.2487	26.906	285	603.006	88.291	26.8667
138	245.819	88.2485	26.903	286	571.366	88.3095	26.8239
139	1051.9	88.2476	26.9052	287	220.043	88.3099	26.8238
140	424.025	88.2481	26.9056	288	500.193	88.3123	26.8269
141	3158.76	88.236	26.9102	289	497.139	88.3259	26.8376
142	1685.25	88.2513	26.908	290	193.821	88.3303	26.8471
143	4283.61	88.2573	26.909	291	25.5029	88.3269	26.8482
144	826.283	88.2398	26.8187	292	16.4798	88.3262	26.8485
145	2661.94	88.2394	26.8289	293	142.254	88.3176	26.8458
146	440.653	88.2386	26.8318	294	54.7755	88.3173	26.8442
147	8140.19	88.2341	26.8994	295	52.0469	88.3134	26.8371
148	803.776	88.2306	26.8658				

Table 2. Elevation character of the Balason river basin and landslide susceptibility

Factors	Subclasses	Total pixels		Landslide occurrence pixels		FR value
		Absolute	Percentage	Absolute	Percentage	
Elevation (meter)	106-298	368681	21.92	4	0.08	0.00
	298-539	139536	8.30	270	4.93	0.59
	539-792	135130	8.03	876	15.96	1.99
	792-1030	172562	10.26	1240	22.60	2.20
	1030-1240	181438	10.79	1457	26.55	2.46
	1240-1440	189984	11.30	573	10.44	0.92
	1440-1650	169253	10.06	273	4.97	0.49
	1650-1880	144806	8.61	75	1.37	0.16
	1880-2130	118740	7.06	180	3.29	0.47
	2130-2603	61869	3.68	539	9.82	2.67

factor, zonal statistics as a table under spatial analyst tool in Arc GIS 10.1 environment was used. In can be mentioned that, higher the frequency ratio value stronger the relationship between landslide occurrence and the given conditioning factor. As indicated by many researchers, natural break method was applied to determine different subclasses of continuous geomorphic parameters (Poli and Sterlacchini, 2007; Althuwaynee et al., 2014; Mahalingam et al., 2016). Finally, the composite index map of landslide susceptibility was obtained using following equation in Arc Map 10.1 GIS environment.

$$CLS = Fr_1 + Fr_2 + Fr_3 + \dots + Fr_n \dots \dots \dots (2)$$

Where, CLS is the composite landslide susceptibility index and $Fr_1 + \dots + Fr_n$ is the frequency ratio of each factor.

RESULTS

Elevation character and landslide susceptibility

The elevation of the Balason river basin was classified into 10 categories where elevation ranges from 106 meter to 2603 meter (Figure 2a). Extreme north-western part and north-eastern part of the basin was dominated by higher elevation whereas, southernmost section was characterized by low elevation. However, no direct relationship can be made between elevation and landslide occurrences. Several researchers showed that landslides have the more tendencies to occur at the higher El areas, so the high El areas have the greater landslide susceptibility (Devkota et al., 2013; Umar et al., 2014). Moderate elevation zones were found in the central portion of the basin. Maximum area of 21.92% was covered by 106-298 meter elevation zone while a small area of 3.68% was dominated by 2130-2603 meter elevation zone. The highest and lowest frequency ratio value was found in the elevation class of 2130-2603 meter (2.67) and 106-298 meter (0.00) respectively, which indicated greater and lower chance of slope failures respectively (Table 2).

Slope angle character and landslide susceptibility

The slope map of the Balason river basin was made from the SRTM DEM and it was classified into 10 different slope zones (Figure 2b). Basically, the river basin is attributed with maximum slope angle where slope angle ranges from 0° to more than 71.10°. The basin is dissected by well developed drainage network which made the slope steepened by continuous branching of the drainage and its headword erosion. The slope plays an important role for the growth and development of drainage network, surface run-off and soil erosion, as well as drainage concentration over the space. Not only that the stream power index (SPI) and topographic wetness index (TWI) are influenced by slope steepness, but the study also revealed that there is a positive relationship between slope and landslide frequency ratio. The slope having greater than 26.60° is attributed with the frequency ratio value ranges from 1.13 to 13.57, which showed high probability of landslide phenomena (Table 3).

Geology and landslide susceptibility

The geology of Balason river basin was divided into five major categories viz., Darjeeling Gneisses, Daling Series, Damuda Series /Lower Gondwanas, Nahan Group and Alluvium (older and recent) (Figure 3a) (Lama, 2003). The Darjeeling gneisses covered a large part of the basin (68.33%) where it varied from a foliated granitoid rock composed to quartz, felsper and biotite to more or less pure mica schist. The Daling series consists of phyllite, slate and quartzite which are being found in the Middle East and Middle West portion of the basin. Damuda Series /Lower Gondwana series consisted of quartzitic sandstones, shales and slates, semi-anthracitic (graphitic) coal, lamprophyre silts and minor bands of limestone. The Nahan Group or Lower Siwalik deposits are mainly composed of soft greyish sandstone, mudstone, shales and conglomerates.

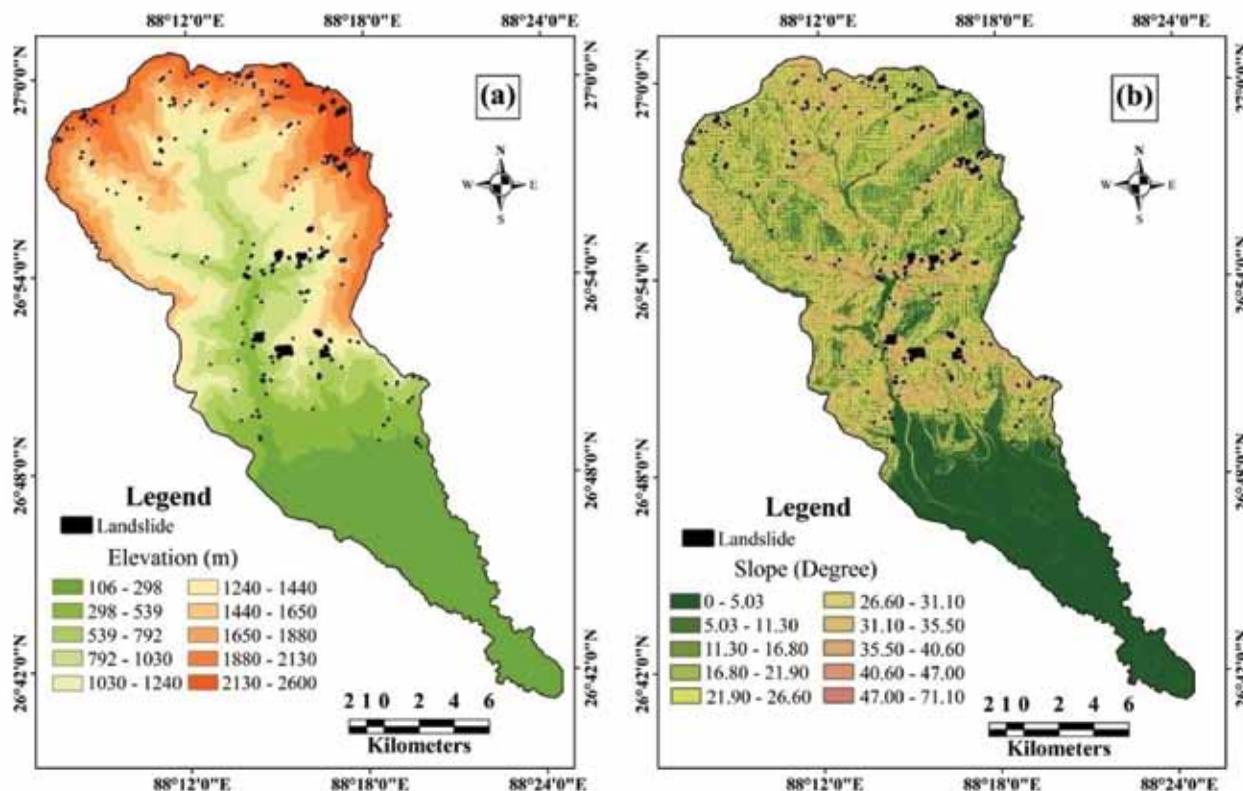


Figure 2. Maps showing data layers of (a) elevation and (b) slope.

Table 3. Slope character of the Balason river basin and landslide susceptibility

Factors	Subclasses	Total pixels		Landslide occurrence pixels		FR value
		Absolute	Percentage	Absolute	Percentage	
Slope (Degree)	0-5.03	400086	23.79	47	0.86	0.04
	5.03-11.30	167825	9.98	202	3.68	0.37
	11.30-16.80	219531	13.05	522	9.52	0.73
	16.80-21.90	231356	13.75	736	13.41	0.98
	21.90-26.60	207903	12.36	601	10.95	0.89
	26.60-31.10	169058	10.05	625	11.38	1.13
	31.10-35.50	126742	7.54	611	11.14	1.48
	35.50-40.60	91415	5.43	663	12.09	2.22
	40.60-47.00	51416	3.06	742	13.52	4.42
47.00-71.10	16667	0.99	738	13.44	13.57	

Alluviums (older and recent) are composed of boulder beds and other sands and gravels. Similarly, Drift Formation and younger flood plain deposits comprise sand and gravel, pebbles, etc. It can be seen that, Darjeeling Gneiss had the greater potentiality of landslide occurrences (1.44) while, Damuda Series /Lower Gondwana and Nahan Group, Lower Siwalik deposits were registered with lowest FR value (0.00), which indicated no chance of slope failures phenomena (Table 4).

Geomorphology and landslide susceptibility

Lama (2003) divided the Balason river basin into 7 distinct geomorphological groups i.e. upper hill (>2000m), middle hill (1000-2000m), lower hill (<1000m), foothill (piedmont), upper alluvial fan, lower alluvial fan and river flood plain (Figure 3b). Middle hill (1000-2000m) covered larger area of the basin (46.22%), while river flood plain had the smallest area of the basin (4.07%). The greater

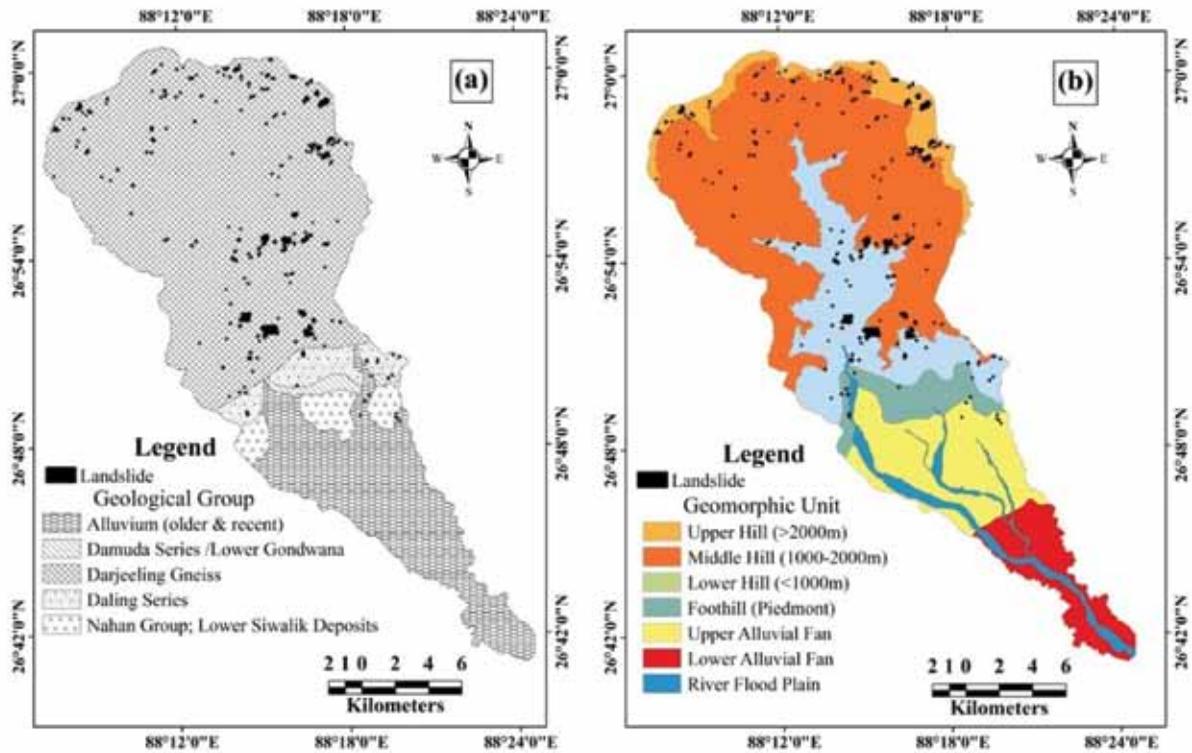


Figure 3. Data layers of (a) geological map and (b) geomorphological map (after Lama, 2003).

Table 4. Geology of the Balason river basin and landslide susceptibility

Factors	Subclasses	Total pixels		Landslide occurrence pixels		FR value
		Absolute	Percentage	Absolute	Percentage	
Geology	Darjeeling Gneiss (Dg)	287309	68.33	1345	98.32	1.44
	Daling Series (Ds)	18518	4.40	21	1.54	0.35
	Damuda Series /Lower Gondwana (DAs)	4784	1.14	0	0.00	0.00
	Nahan Group, Lower Siwalik Deposits (Ng)	16839	4.00	0	0.00	0.00
	Alluvium (older & recent) (Al)	93050	22.13	2	0.15	0.01

potentiality of landslide occurrences was found in lower hill (<2000m) (1.90). The chance of slope failures was lowest in the class of lower alluvial fan and river flood plain, both were registered with 0.00 FR value (Table 5).

Soil character and landslide susceptibility

Five types of soil were identified in the basin i.e. W002, W003, W004, W006 and W009 (Figure 4a). The soil of the Balason river basin is favorable for tea plantation. The original soil of the area is drastically changed either by the process of new soil formation on the truncated top or by removal of soil by erosion (Sarkar, 1990). The parent soil of this region has changed due to its nature and it has formed more than one time (Lama, 2003). In the sandy

area, soils are generally siliceous and aluminous. The lower Balason basin is formed by fine grained clay loam. In some cases, the top soils are finer in nature than the subsurface soils. The detailed characteristics of each type of soil of the Balason river basin were presented in Table 7. In this basin, W003 soil type was registered with highest frequency ratio is 1.88. On the other hand, W006 and W009 soil types covering almost 17% area of the basin, have the lowest chance of landslide events (Table 8).

Spatial distribution of Rainfall and landslide susceptibility

Landslides are closely associated with duration, amount and intensity of rainfall. The intensity and amount of

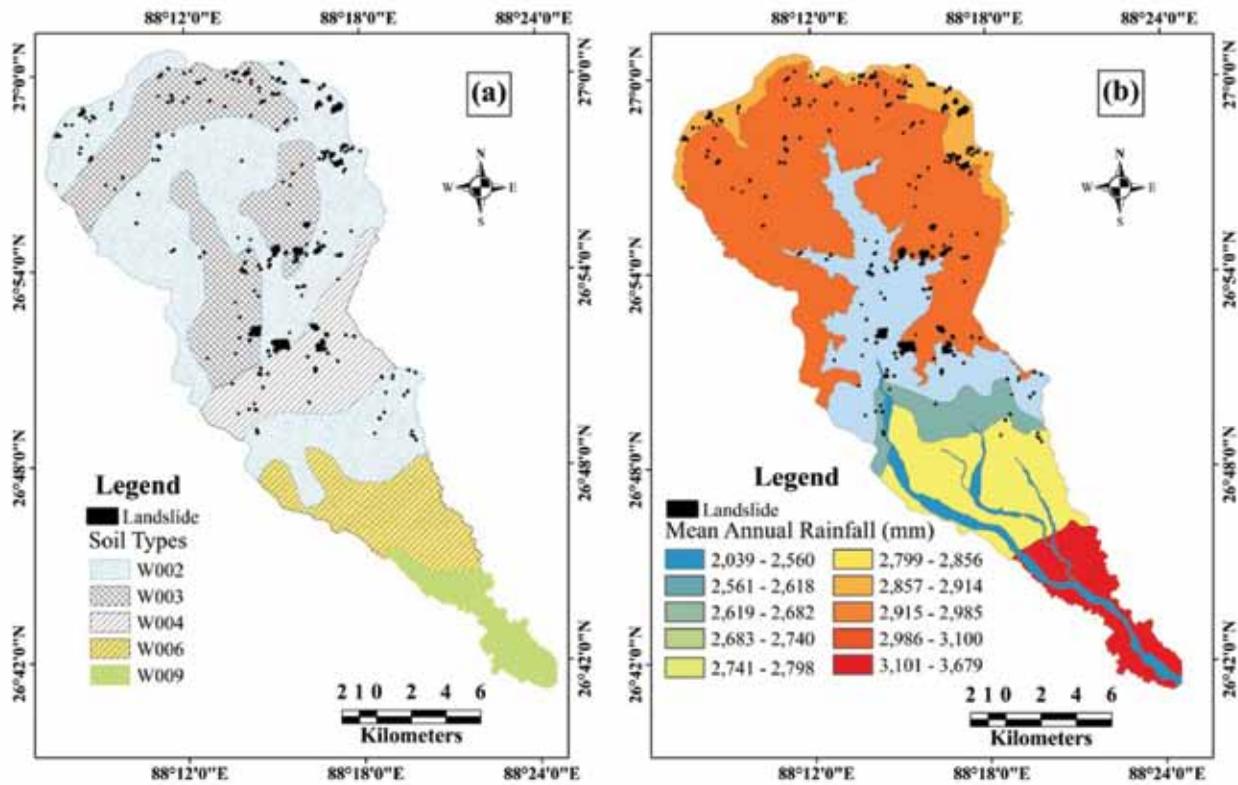


Figure 4. Data layers of (a) soil and (b) rainfall maps.

Table 5. Geomorphology of the Balason river basin and landslide susceptibility

Factors	Subclasses	Total pixels		Landslide occurrence pixels		FR value
		Absolute	Percentage	Absolute	Percentage	
Geomorphology	Upper Hill (>2000m)	29309	6.97	150	10.96	1.57
	Middle Hill (1000-2000m)	194339	46.22	732	53.51	1.16
	Lower Hill (<2000m)	77058	18.33	477	34.87	1.90
	Foothill (Piedmont)	21111	5.02	7	0.51	0.10
	Upper Alluvial Fan	53984	12.84	2	0.15	0.01
	Lower Alluvial Fan	27588	6.56	0	0.00	0.00
	River Flood Plain	17111	4.07	0	0.00	0.00

rainfall varies from one place to another in the basin due to topographic configuration. The mean annual rainfall ranges from 2041 mm to 3673 mm. Based on the rainfall distribution the Balason river basin was classified into ten zones (Figure 4b). Around 35% of the landslide-affected area and 50% of the basin are characterized by mean annual rainfall of more than 2790 mm. More than 1 frequency ratio was found in the class ranges from 2559 to 2847 mm and 3090-3673 mm rainfall, which was characterized by high landslide occurrence events. Rest of the class having less than 1 frequency ratio value, which indicates a minimum chance of landslide occurrences. Some high annual rainfall classes, i.e. 2847-2905 mm, 2905-2975 mm, and 2975-3090 mm, have low FR values because of the unfavorable geo-

environmental conditions for landslide occurrences in the concerned area (Table 9).

Drainage density and landslide susceptibility

The drainage network development on both sides of the slope segments makes the interfluvial area more steep, resulting in mountain slope instability. The drainage network development, slope steepening, and slope failure are regular processes which promote equilibrium in the geomorphic system. The concentration of the drainage network and their engagement in the process of erosion and transportation make the slope more vulnerable to landslide phenomena. The drainage density varies from one place to another based

Table 7. Soil characteristics of Balason river basin

Symbol	Soil characteristics	Taxonomic name
W002	Moderately shallow, excessively drained, coarse loamy soils occurring on steep side slopes with gravelly loamy surface, severe erosion and strong rockiness, associated with moderately shallow, well drained, gravelly loamy soils with loamy surface and moderate erosion.	Coarse loamy, typic Udorthents; Loamy skeletal, typic Dystrochrepts
W003	Deep, well drained, fine loamy soils occurring on steep side slopes with gravelly loamy surface, moderate erosion and moderate rockiness, associated with moderately shallow, excessively drained, coarse loamy soils with loamy surface, severe erosion and moderate rockiness.	Fine loamy, Umbric Dystrochrepts; Coarse loamy, Typic Udorthents
W004	Moderately shallow, well drained, gravelly loamy soils occurring on steep side slopes with gravelly loamy surface, moderate erosion and moderate rockiness, associated with moderately shallow, somewhat excessively drained, gravelly loamy soil with loamy surface, moderate erosion and moderate rockiness	Loamy skeletal, Typic Haplumbrepts; Loamy skeletal, typic Udorthents
W006	Very deep, imperfectly drained, coarse loamy soils occurring on very gently sloping upper piedmont plains with loamy surface and moderate erosion, associated with very deep, imperfectly drained fine loamy soils.	Coarse loamy, Umbric Dystrochrepts; Fine loamy, Fluventic Dystrochrepts
W009	Very deep, imperfectly drained, coarse loamy soils occurring on nearly level lower piedmont plain with loamy surface, associated with very deep imperfectly drained, fine loamy soils.	Coarse loamy, Aquic Udifluvents; Fine loamy, Fluventic Eutrochrepts

Table 8. Soil of the Balason river basin and landslide susceptibility.

Factors	Subclasses	Total pixels		Landslide occurrence pixels		FR value
		Absolute	Percentage	Absolute	Percentage	
Soil	W002	198569	47.22	524	38.30	0.81
	W003	102249	24.32	626	45.76	1.88
	W004	49222	11.71	218	15.94	1.36
	W006	46327	11.02	0	0.00	0.00
	W009	24133	5.74	0	0.00	0.00

Table 9. Rainfall of the Balason river basin and landslide susceptibility

Factors	Subclasses	Total pixels		Landslide occurrence pixels		FR value
		Absolute	Percentage	Absolute	Percentage	
Rainfall (mm)	2041-2559	40635	9.66	130	9.48	0.98
	2559-2617	42137	10.02	209	15.25	1.52
	2617-2681	44545	10.59	186	13.58	1.28
	2681-2732	38867	9.24	161	11.80	1.28
	2732-2790	45248	10.76	186	13.58	1.26
	2790-2847	41930	9.97	147	10.72	1.08
	2847-2905	41695	9.92	87	6.36	0.64
	2905-2975	42379	10.08	40	2.91	0.29
	2975-3090	42956	10.22	62	4.53	0.44
	3090-3673	40108	9.54	161	11.80	1.24

on the developed drainage network in the basin. Higher the drainage density more is the drainage concentration and slope saturation over the space. The drainage density map was divided into 10 classes (Figure 5a). It was observed that moderate level of drainage density have high frequency ratio

and greater probability of slope instability. High drainage density of more than 5.39 cover about 21% area of the basin. 10.60-15.10 class have maximum chance of landslide occurrences (2.94), whereas 0-1.31 drainage density class have least chance of slope failures (0.00) (Table 10).

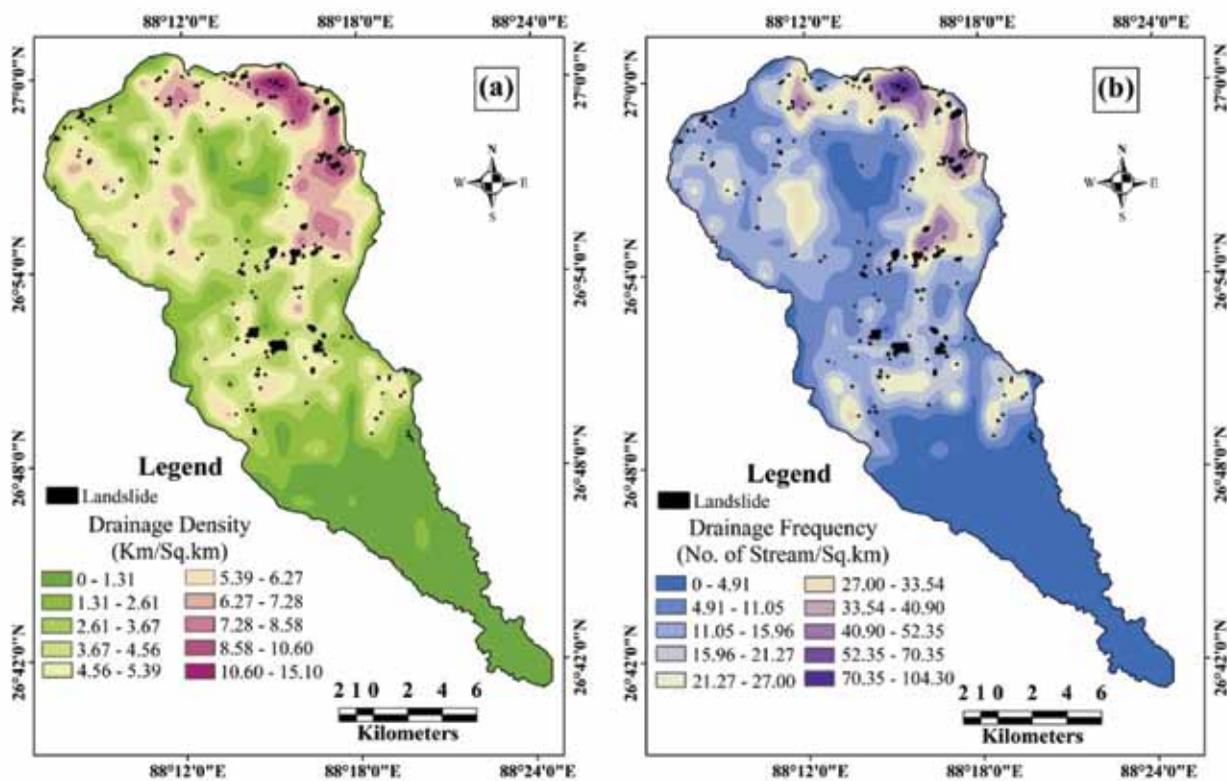


Figure 5. Data layers of (a) drainage density map and (b) drainage frequency map.

Table 10. Drainage density of the Balason river basin and landslide susceptibility

Factors	Subclasses	Total pixels		Landslide occurrence pixels		FR value
		Absolute	Percentage	Absolute	Percentage	
Drainage density (Km/Sq. Km)	0-1.31	338135	20.10	4	0.07	0.00
	1.31-2.61	197209	11.72	8	0.14	0.01
	2.61-3.67	239068	14.21	783	14.27	1.00
	3.67-4.56	284654	16.92	1830	33.35	1.97
	4.56-5.39	246882	14.68	1108	20.19	1.38
	5.39-6.27	177847	10.57	746	13.59	1.29
	6.27-7.28	103596	6.16	494	9.00	1.46
	7.28-8.58	60735	3.61	366	6.67	1.85
	8.58-10.60	25924	1.54	73	1.33	0.87
	10.60-15.10	7951	0.47	76	1.39	2.94

Drainage frequency and landslide susceptibility

Similar to the drainage density map, drainage frequency of the basin was also divided into 10 classes (Figure 5b). About 67% area of the basin registered drainage frequency value of 0 to 15.96. 70.35-104.30 class has the greater potentiality in landslide occurrences (4.59) and 0-4.91 class have the minimum probability of slope failure events (0.00) (Table 11). In addition, the classes having drainage frequency value of 33.54-40.90 (4.37), 52.35-70.35 (2.76),

11.05-15.96 (2.17) and 27.00-33.54 (1.87) were also characterized by high landslide susceptibility.

DISCUSSION

The trend line of FR values reveals that elevation, slope, drainage density and drainage frequency were positively correlated with slope failure which means that slope instability increasing with increasing the value (Figure 6a, b, d and e). On the other hand, there was negative relation

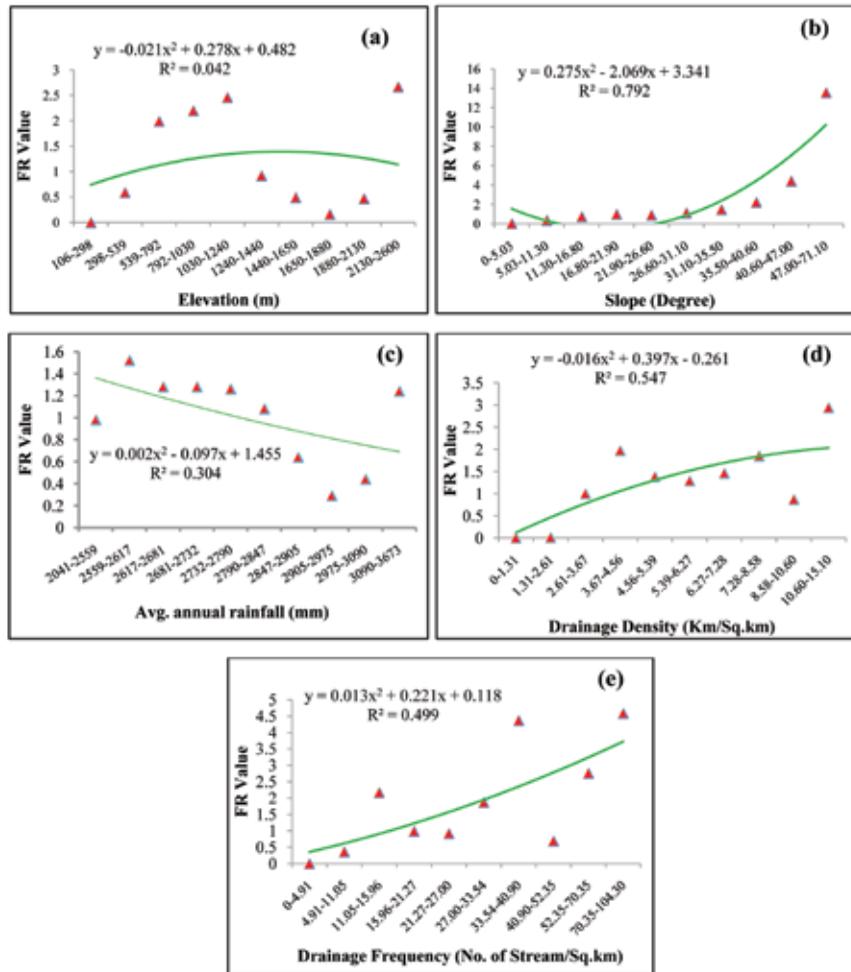


Figure 6. Relation between data layers and FR value (a) elevation (b) slope (c) rainfall (d) drainage density and (e) drainage frequency.

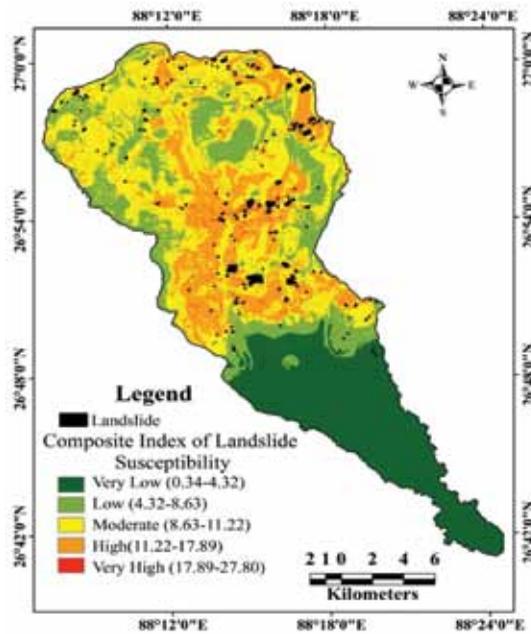


Figure 7. Composite index map of landslide susceptibility using FR value.

Table 11. Drainage frequency of the Balason river basin and landslide susceptibility

Factors	Subclasses	Total pixels		Landslide occurrence pixels		FR value
		Absolute	Percentage	Absolute	Percentage	
Drainage frequency (No. of stream/Sq.km)	0-4.91	446653	26.55	4	0.07	0.00
	4.91-11.05	315499	18.76	370	6.74	0.36
	11.05-15.96	374903	22.29	2657	48.42	2.17
	15.96-21.27	260262	15.47	842	15.35	0.99
	21.27-27.00	131533	7.82	394	7.19	0.92
	27.00-33.54	73153	4.35	446	8.13	1.87
	33.54-40.90	42087	2.50	600	10.94	4.37
	40.90-52.35	27831	1.65	62	1.14	0.69
	52.35-70.35	6621	0.39	60	1.09	2.76
70.35-104.30	3458	0.21	52	0.94	4.59	

Table 12. Spatial correlation between the landslide susceptibility index values and the various geomorphic parameters

Model/ geomorphic parameters	Elevation	Slope	Geology	Geomorphology	Soil	Rainfall	Drainage density	Drainage frequency
Landslide susceptibility index values	0.67	0.62	0.79	0.83	0.68	0.36	0.72	0.61

between rainfall and landslide occurrences (Figure 6c). In case of geology, soil and geomorphology, trend line of FR value was not shown in a diagram due to its categorical nature. In addition, geomorphology was most intimately associated with the determination of landslide susceptibility with correlation value of 0.83. Next highest relationship with landslide susceptibility was observed in case of geology (0.79) followed by drainage density (0.72), elevation (0.67), slope (0.62), drainage frequency (0.61) respectively. Lastly, rainfall exerted lowest influence on landslide susceptibility as indicated by correlation value of 0.36 (Table 12). Finally, composite FR index of landslide susceptibility was made by considering the eight factors, and divided it into five landslides susceptibility zones using natural breaks classification method i.e. very low (0.34-4.32), low (4.32-8.63), moderate (8.63-11.22), high (11.22-17.89) and very high (17.89-27.80) (Figure 7).

CONCLUSIONS

All the geomorphic parameters play a significant role in assessing landslide susceptibility in the Balason basin. Not only the factors but also the individual category of all the landslide occurrence factors showed a distinct level of landslide probability. Generally elevation, slope, drainage density and drainage frequency were positively correlated with landslide susceptibility, whereas rainfall was negatively associated with landslide occurrences. It is assumed that the prepared landslide susceptibility analysis will help the policy makers, developers and engineers for choosing suitable

locations to implement slope management plans, land use plans and other development action plans, although this method can be less useful on the site-specific scale.

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Compliance with Ethical Standards

The authors declare that they have no conflict of interest and adhere to copyright norms.

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